

A Statistical Mechanics approach to Geophysical Turbulence

Corentin Herbert
ENS de Lyon

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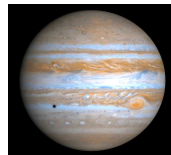
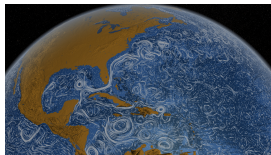
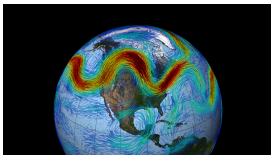
In collaboration with:

- B. Dubrulle* SPEC, CEA Saclay, France
- G. Falkovich* Weizmann Institute, Israel
- A. Frishman* Princeton University, USA
- A. Pouquet* NCAR, Boulder, USA
- D. Rosenberg* ORNL, Oak Ridge, USA & SciTec., Princeton, USA
- F. Bouchet* ENS Lyon, France



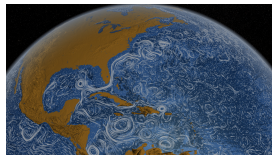
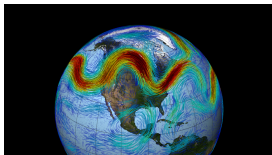
Motivation: geophysical flows

Geophysical flows self-organize into large-scale coherent structures, which play a major part in weather/climate.

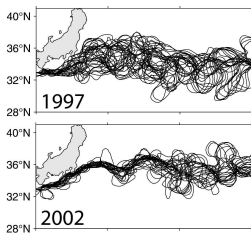


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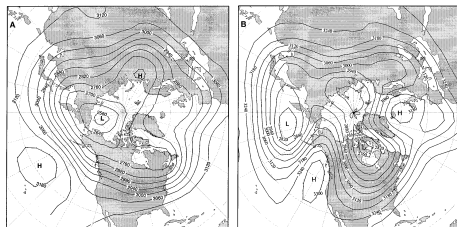
Geophysical flows self-organize into large-scale coherent structures, which play a major part in weather/climate.



They fluctuate and undergo abrupt transitions.



Kuroshio path¹



Zonal/blocked Jet Stream transition²

¹B Qiu and S. M. Chen (2005).. *J. Phys. Oceanogr.*

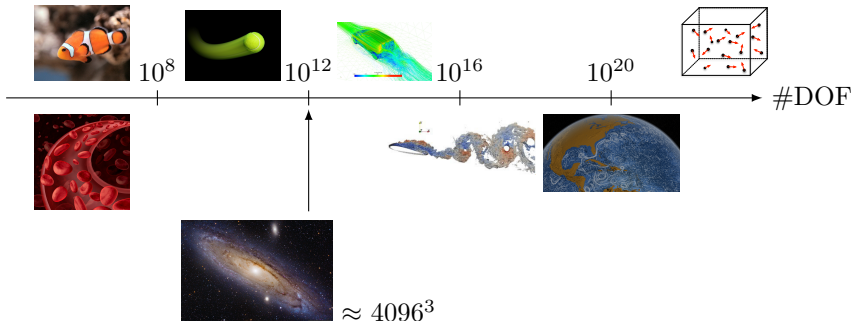
²E. R. Weeks et al. (1997).. *Science*

Turbulent flows and degrees of freedom

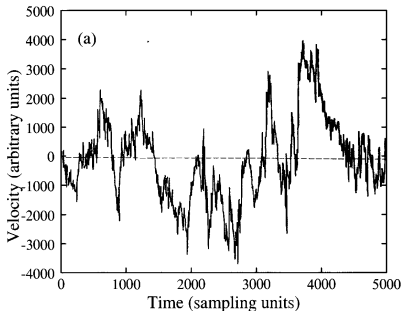
Navier-Stokes equations:

$$\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla P + \nu \Delta \mathbf{u}, \quad \text{Re} = UL/\nu$$

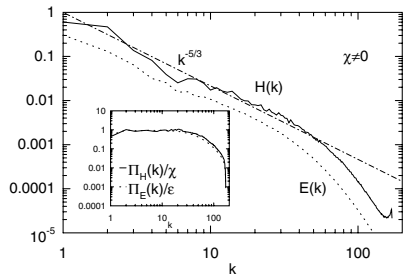
- ▶ Chaotic nature: nonlinear term couples wide range of scales.
- ▶ Number of degrees of freedom $\sim \text{Re}^{9/4}$.



Predictable and unpredictable observables



Velocity measurement; ONERA wind tunnel (Y. Gagne, E. Hopfinger)



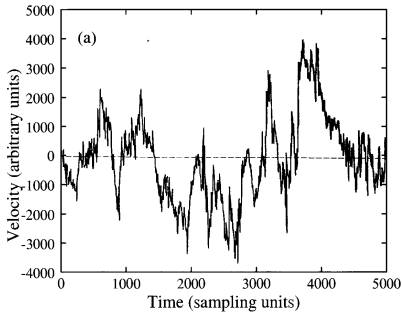
Universal kinetic energy spectrum in a DNS³.

Can we find a probability distribution describing the system?

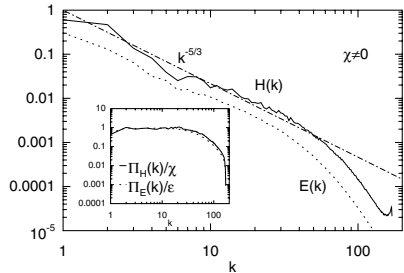
Very difficult task!

³Q. Chen et al. (2003).. *Phys. Rev. Lett.*

Predictable and unpredictable observables



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Universal kinetic energy spectrum in a DNS³.

Kolmogorov theory Probability distribution respects symmetries (homogeneity, isotropy, scale invariance)

Not true: e.g. scale invariance is spontaneously broken (*intermittency*).

Geophysical flows break the symmetries of classical turbulence, which allows for new theoretical approaches.

³Q. Chen et al. (2003).. *Phys. Rev. Lett.*

Dynamical Models: 2D and Geophysical Turbulence

Model for incompressible turbulent flows (*Navier-Stokes equation*):

$$\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla P + \nu \Delta \mathbf{u} + \mathbf{F},$$
$$\nabla \cdot \mathbf{u} = 0.$$

When $\nu = F = 0$ (no forcing and no dissipation), we have the *Euler equations*.

In terms of *vorticity* $\boldsymbol{\omega} = \nabla \times \mathbf{u}$,

- ▶ **For a 2D domain**, $\boldsymbol{\omega} = \omega \mathbf{n}$, vorticity is conserved along trajectories (*Lagrangian invariant*):

$$\partial_t \omega + \mathbf{u} \cdot \nabla \omega = 0.$$

- ▶ **For a 3D domain**, it is not:

$$\partial_t \boldsymbol{\omega} + \mathbf{u} \cdot \nabla \boldsymbol{\omega} = \boldsymbol{\omega} \cdot \nabla \mathbf{u}.$$

Dynamical Models: 2D and Geophysical Turbulence

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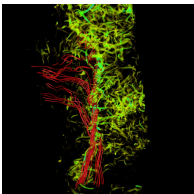
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Radically different behavior:

3D HIT

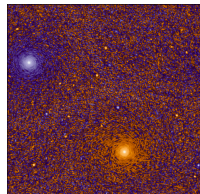
Fig. A. Pouquet (NCAR)



Direct cascade
Intermittency

...

2D Turbulence



Inverse cascade
Conformal invariance

...

Dynamical Models: 2D and Geophysical Turbulence

Model for incompressible turbulent flows (*Navier-Stokes equation*):

$$\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla P + \nu \Delta \mathbf{u} + \mathbf{F},$$
$$\nabla \cdot \mathbf{u} = 0.$$

When $\nu = F = 0$ (no forcing and no dissipation), we have the *Euler equations*.

Invariants:

3D HIT

$$\text{Energy } E = \frac{1}{2} \int \mathbf{u}^2(\mathbf{r}) d\mathbf{r}$$

2D Turbulence

$$\text{Energy } E = \frac{1}{2} \int \omega(\mathbf{r}) \psi(\mathbf{r}) d\mathbf{r} \quad (\omega = -\Delta \psi)$$

$$\text{Casimir invariants } \int s(\omega(\mathbf{r})) d\mathbf{r}$$

$$\text{E.g. } \textit{enstrophy} \int \omega^2(\mathbf{r}) d\mathbf{r}$$

Dynamical Models: 2D and Geophysical Turbulence

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When $\nu = F = 0$ (no forcing and no dissipation), we have the *Euler equations*.

- ▶ **For a 2D domain**, $\omega = \omega \mathbf{n}$, vorticity is conserved along trajectories (*Lagrangian invariant*):

$$\partial_t \omega + \mathbf{u} \cdot \nabla \omega = 0.$$

- ▶ **Geophysical flows** are 3D, but subjected to strong rotation and density stratification.

Large scales well described by advection of *potential vorticity* (*quasi-geostrophic*):

$$\partial_t q + \mathbf{u} \cdot \nabla q = 0.$$

E.g. $q = \omega + \partial_z(f_0^2/N^2 \partial_z \psi) + \beta y$.

Main Questions and Theoretical Tools

Generic questions:

- ▶ Can we predict self-organization of geophysical flows into large scale coherent structures?
- ▶ Characterize the attractors of geophysical turbulence
- ▶ Study fluctuations around the mean state
- ▶ What aspects of transitions in turbulent flows are predictable?

Because of strong nonlinearity/huge number of degrees of freedom, classical fluid mechanics+direct numerical simulations do not suffice.

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Main theoretical tool: Large Deviation Theory

It is a tool to study asymptotic probabilities:

$$\text{Prob}(A[x] = a) \sim e^{-I(a)/\varepsilon} \text{ when } \varepsilon \rightarrow 0.$$

The small parameter ε can be

- ▶ The inverse of the number of degrees of freedom $\varepsilon = 1/N$.
- ▶ The amplitude of a noise term (Freidlin-Wentzell theory)
- ▶ The inverse of an observation time $\varepsilon = 1/T$ (Donsker-Varadhan)

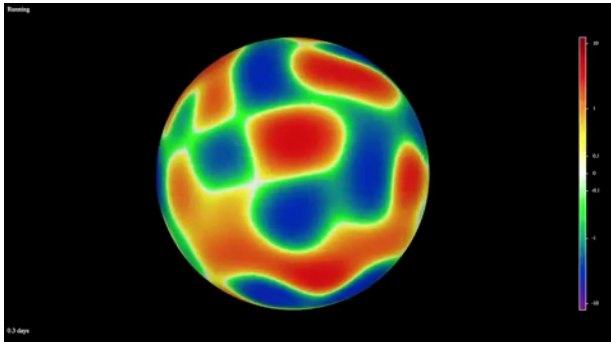
Outline

- 1 Introduction
- 2 Equilibrium Theory**
- 3 Perturbative approach
- 4 Large deviations and transitions
- 5 Conclusion

Advection of the vorticity field in 2D

$$\partial_t \omega + \mathbf{u} \cdot \nabla \omega = 0.$$

Small-scale vorticity is mixed by the flow while large-scale coherent structures form.



Direct Numerical Simulation: Vorticity Contours. Courtesy Brad Marston (Brown University).

The microcanonical measure

Formally, we define the *microcanonical measure* as

$$\mu_{E,(\Gamma_n)_{n \in \mathbb{N}}}(d\omega) = \frac{1}{\Omega(E, (\Gamma_n)_{n \in \mathbb{N}})} \delta(\mathcal{E}[\omega] - E) \prod_{k=1}^{+\infty} \delta(\mathcal{G}_k[\omega] - \Gamma_k) \prod_{i=1}^{+\infty} d\omega_i.$$

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- ▶ Invariant measure of the Euler equations⁴.
- ▶ Difficult to manipulate: e.g.

$$\mathbb{E}[\omega] = 0.$$

Spontaneous symmetry breaking.

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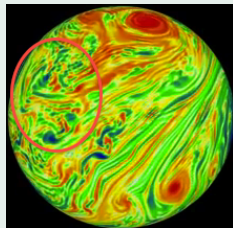
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Can we compute macrostates?

Mean-field theory (Miller-Robert-Sommeria)

Two levels of description⁴:

- ▶ **Microstates**: fine-grained vorticity field $\omega(\mathbf{x})$.



⁴R. Robert and J. Sommeria (1991).. *J. Fluid Mech.* J. Miller (1990).. *Phys. Rev. Lett.*

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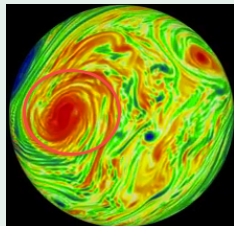
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Can we compute macrostates?

Mean-field theory (Miller-Robert-Sommeria)

Two levels of description⁴:

- ▶ Microstates: fine-grained vorticity field $\omega(\mathbf{x})$.
- ▶ **Macrostates**: fine-grained vorticity probability distribution $\rho(\sigma, \mathbf{x})$,
 $\int \rho(\sigma, \mathbf{x}) d\sigma = 1$.
 Mean coarse-grained vorticity:
 $\bar{\omega}(\mathbf{x}) = \int \sigma \rho(\sigma, \mathbf{x}) d\sigma$.

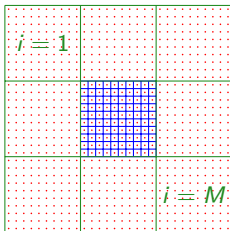


We want to compute the most probable macrostates ρ

⁴R. Robert and J. Sommeria (1991).. *J. Fluid Mech.* J. Miller (1990).. *Phys. Rev. Lett.*

The mean-field approach: counting the microstates⁵

Let us consider a square lattice with N sites, and a “coarse-grained” lattice of M boxes containing $n = N/M$ sites each.



Finite number of vorticity levels $\mathfrak{S} = \{\sigma_1, \dots, \sigma_K\}$.

► *Microstates:*

$$\hat{\omega} = (\omega_{i\alpha})_{\substack{1 \leq i \leq M \\ 1 \leq \alpha \leq n}} \in \mathfrak{S}^N.$$

► *Macrostates:*

$$P = (p_{ik})_{\substack{1 \leq i \leq M \\ 1 \leq k \leq K}} \in [0, 1]^{MK}, \quad \sum_{k=1}^K p_{ik} = 1.$$

► *Coarse-grained vorticity field:*

$$\bar{\omega}_i \equiv \frac{1}{n} \sum_{\alpha=1}^n \omega_{i\alpha} = \sum_{k=1}^K \sigma_k p_{ik}.$$

Number of microstates which realize a given macrostate:

$$W(P) = \prod_{i=1}^M \frac{n!}{\prod_{k=1}^K (np_{ik})!}$$

⁵C. Herbert (2015).. In: *Stochastic Equations for Complex Systems: Theoretical and Computational Topics*. Ed. by S. Heinz and H. Bessaih. Springer

The mean-field approach: large deviation of the macrostate probability⁶

Conservation constraints:

- ▶ *Vorticity distribution* (i.e. Casimir invariants) depends only on P
- ▶ *Energy* depends only on P in the limit $N \rightarrow +\infty$.

Probability of a given macrostate P with energy E :

$$\text{Prob}(P) = \frac{W(P)}{\Omega_N(E, \gamma)},$$
$$\frac{1}{N} \ln \text{Prob}(P) = -\frac{1}{M} \underbrace{\sum_{i=1}^M \sum_{k=1}^K p_{ik} \ln p_{ik}}_{\text{entropy } \mathcal{S}_{M,K}[P]} - S(E, \gamma) + o(1).$$

This is a *large deviation property*.

The mean-field approach: variational problem⁷

Equilibrium states = most probable macrostates. They must minimize the large deviation rate function, while satisfying the global constraints.

Microcanonical variational problem

$$S(E, \gamma) = \max_{\rho} \{ \mathcal{S}[\rho] \mid \mathcal{E}[\rho] = E, \forall \sigma \in \mathbb{R}, \mathcal{D}_{\sigma}[\rho] = \gamma(\sigma) \}.$$

Critical points:

$$\rho(\sigma, \mathbf{r}) = \frac{e^{-\beta\sigma\bar{\psi}(\mathbf{r}) - \alpha(\sigma)}}{\mathcal{Z}_{\beta, \alpha}(\bar{\psi}(\mathbf{r}))} \quad (\text{Gibbs states}),$$

with

$$\bar{\omega} \equiv -\Delta\bar{\psi}, \quad \mathcal{Z}_{\beta, \alpha}(u) \equiv \int_{\mathbb{R}} e^{-\beta\sigma u - \alpha(\sigma)} d\sigma.$$

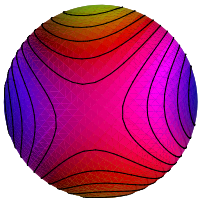
Mean-field equation:

$$\bar{\omega}(\mathbf{r}) = F_{\beta, \alpha}(\bar{\psi}(\mathbf{r})), \quad \text{with } F_{\beta, \alpha}(u) \equiv -\frac{1}{\beta} \frac{d \ln \mathcal{Z}_{\beta, \alpha}(u)}{du}.$$

Exemple: Equilibrium flows on the sphere

Stable equilibrium states⁸

- ▶ Solid body rotations: $\psi = \Omega_* \cos \theta$
- ▶ Dipoles: $\psi = \Omega_* \cos \theta + \sqrt{3(E - E^*(L))} \sin \theta \cos(\phi - \phi_0)$
- ▶ *Quadrupoles:*
$$\psi_\infty = \psi_{20}(3 \cos^2 \theta - 1) + \psi_{21} \sin(2\theta) \sin(\phi - \phi_1) + \psi_{22} \sin^2 \theta \sin(2(\phi - \phi_2))$$



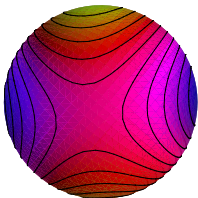
Theoretical Equilibrium: Quadrupole

⁸C. Herbert et al. (2012).. *J. Stat. Mech.* C. Herbert (2013).. *J. Stat. Phys.*

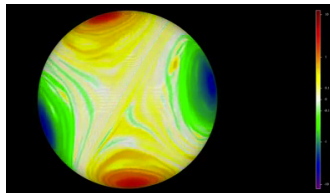
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DNS Final State⁹

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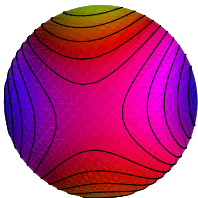
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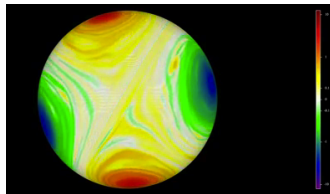
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Theoretical Equilibrium: Quadrupole



DNS Final State⁹

Generalization to more realistic geophysical flows¹⁰

⁸C. Herbert et al. (2012).. *J. Stat. Mech.* C. Herbert (2013).. *J. Stat. Phys.*

⁹W. Qi and J. B. Marston (2014).. *J. Stat. Mech.*

¹⁰F. Bouchet and A. Venaille (2012).. *Phys. Rep.* C. Herbert (2014).. *Phys. Rev. E*; V. Lucarini et al. (2014).. *Rev. Geophys.* A. Renaud et al. (2016).. *J. Stat. Phys.*

Summary

Achievements

- ▶ The Microcanonical measure can be built without UV divergences.
- ▶ Mean-field theory is exact: in the microcanonical ensemble, vorticity at two different points behaves as statistically independent random variables.
- ▶ Macrostates satisfy a large deviation property. Equilibrium states can be computed as solutions of a variational problem.
- ▶ They are in qualitative agreement with stationary state of numerical simulations.
- ▶ *(Interesting thermodynamical properties (long-range interactions): non-equivalence of ensembles, negative temperatures, etc)*

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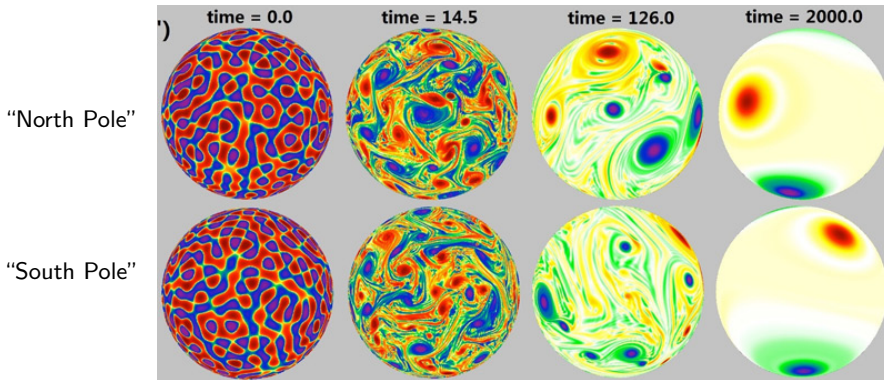
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Limitations

- ▶ Non-ergodicity

The effect of rotation (DNS results¹¹)

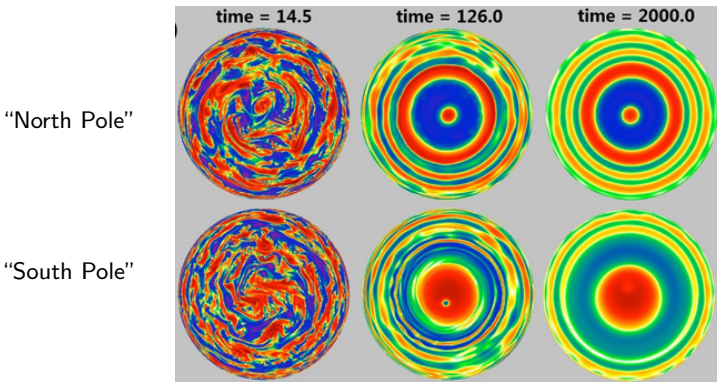
When the Rossby waves are sufficiently slow, the system relaxes towards its equilibrium state.



¹¹W. Qi and J. B. Marston (2014).. *J. Stat. Mech.*

The effect of rotation (DNS results¹¹)

For faster rotation rates, Rossby waves arrest the cascade at the Rhines scale and lead to the emergence of zonal flows.



¹¹W. Qi and J. B. Marston (2014).. *J. Stat. Mech.*

Summary

Achievements

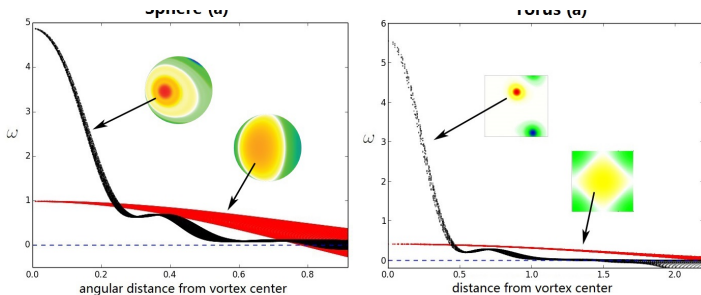
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Limitations

- ▶ Non-ergodicity
- ▶ **Quantitative predictions are difficult. The set of MRS equilibria is huge.**

Vortex profile and higher-order invariants

Comparing the equilibrium states with numerical simulations¹²:



Perturbative expansion leads to core sharpening, but it is difficult to make quantitative predictions.

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¹²W. Qi and J. B. Marston (2014).. *J. Stat. Mech.*

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Limitations

- ▶ Non-ergodicity
- ▶ Quantitative predictions are difficult. The set of MRS equilibria is huge.
- ▶ **Forcing and dissipation not taken into account**

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- 2 Equilibrium Theory
- 3 Perturbative approach**
- 4 Large deviations and transitions
- 5 Conclusion

The closure problem for Homogeneous Isotropic Turbulence

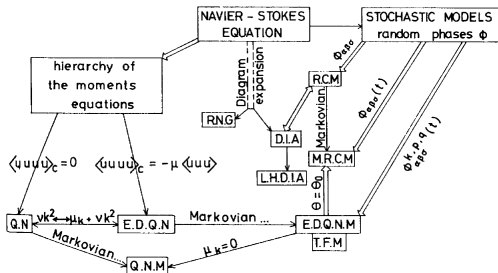
Incompressible Navier-Stokes equation in Fourier space:

$$(\partial_t + \nu k^2)\hat{u}_i(\mathbf{k}) = \sum_{\mathbf{p}, \mathbf{q}} \mathcal{P}_i^{jl}(\mathbf{k})\delta(\mathbf{k} - \mathbf{p} - \mathbf{q})\hat{u}_j(\mathbf{p})\hat{u}_l(\mathbf{q}), \quad k^i \hat{u}_i(\mathbf{k}) = 0.$$

Formally,

$$\begin{aligned} [\partial_t + \nu(k^2 + p^2)]\langle \hat{u}(\mathbf{k})\hat{u}(\mathbf{p}) \rangle &= \langle \hat{u} \star \hat{u} \star \hat{u} \rangle, \\ [\partial_t + \nu(k^2 + p^2 + q^2)]\langle \hat{u}(\mathbf{k})\hat{u}(\mathbf{p})\hat{u}(\mathbf{q}) \rangle &= \langle \hat{u} \star \hat{u} \star \hat{u} \star \hat{u} \rangle, \\ &\dots \end{aligned}$$

Closing the hierarchy requires arbitrary hypothesis (e.g. Gaussianity, etc)



Analytical closure theories¹³

The closure problem for Homogeneous Isotropic Turbulence

Reynolds decomposition:

$$u_i = \bar{u}_i + u'_i,$$

where $\bar{\cdot}$ is a projection operator.

The Navier-Stokes equations become:

$$\begin{aligned}\partial_t \bar{u}_i + \bar{u}_j \partial^j \bar{u}_i &= -\partial_i \bar{P} + \nu \partial_j \partial^j \bar{u}_i - \partial^j \overline{u'_i u'_j}, \\ \partial_t u'_i + \bar{u}_j \partial^j u'_i + u'_j \partial^j \bar{u}_i &= -\partial_i P' + \nu \partial_j \partial^j u'_i - \partial^j u'_i u'_j + \partial^j \overline{u'_i u'_j}.\end{aligned}$$

Modeling approaches:

- ▶ *Large Eddy Simulations*: spatial filtering

$$\bar{u}_i(\mathbf{x}, t) = \int G(\mathbf{x} - \mathbf{y}) u_i(\mathbf{y}, t) d\mathbf{y}$$

- ▶ *Reynolds Average Navier-Stokes*: time filtering

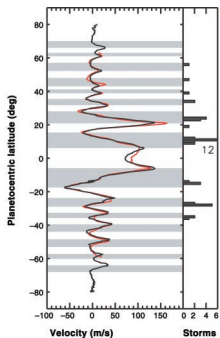
These are *phenomenological models*.

The major difficulty is to compute the Reynolds stress tensor $-\partial^j \overline{u'_i u'_j}$.

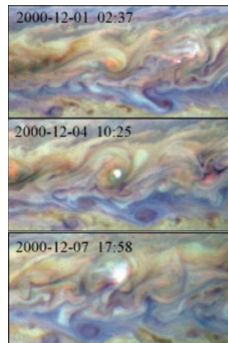
Timescale separation in geophysical flows

In some flows, there is a *natural timescale separation*, usually associated to a broken symmetry of the Navier-Stokes equations.

E.g. Jupiter¹⁴:



Zonal wind measured by Voyager 2 (1979, red) and Cassini (2000, black).



Cassini

¹⁴C. C. Porco et al. (2003).. *Science*.

Adiabatic elimination of fast variables¹⁵ (*stochastic averaging*)

Slow-fast SDE:

$$dX_t = f(X_t, Y_t)dt + \sqrt{2\epsilon}dW_t,$$

$$dY_t = \alpha^{-1}g(X_t, Y_t)dt + \sqrt{\alpha^{-1}}h(X_t, Y_t)dW_t.$$

- ▶ Joint PDF $P(x, y; t)$; Fokker-Planck equation $\partial_t P = (\alpha^{-1}L_0 + L_1)P$.
- ▶ Stationary distribution for fast modes at fixed x and projection operator:

$$L_0 P_\infty^x(y) = 0, \quad \mathcal{P}\phi = P_\infty^x(y) \int dy \phi(x, y).$$

- ▶ Write $P_s = \mathcal{P}P$, $P_f = (1 - \mathcal{P})P$. We have
 $\partial_t P_s = \mathcal{P}(\alpha^{-1}L_0 + L_1)P = \mathcal{P}L_1 P$.
- ▶ At lowest order, $\partial_t P_s = \mathcal{P}L_1 P_s + O(\alpha)$ and $P_s(x, y) = P_\infty^x(y)Q(x)$ with

$$\frac{\partial Q}{\partial t} = -\frac{\partial}{\partial x} [\mathbb{E}_\infty^x[f]Q(x)] + \epsilon \frac{\partial^2}{\partial x^2} Q + O(\alpha).$$

Finally, after adiabatic reduction:

$$dX_t = \mathbb{E}_\infty^{X_t}[f]dt + \sqrt{2\epsilon}dW_t.$$

¹⁵e.g. C. W. Gardiner (2009). *Handbook of Stochastic Methods for physics, chemistry, and the natural sciences*. 4th edition. Springer, Berlin.

Adiabatic elimination of fast variables: zonal jets

Reynolds decomposition for the zonal jets

$\omega = \bar{\omega} + \omega'$, with $\bar{\cdot}$ the projection on the (slow) zonal modes.
Formally,

$$\begin{aligned}\partial_t \bar{\omega} + L_{\bar{\omega}}[\bar{\omega}] &= -\partial^i \overline{u'_i \omega'} + \bar{\eta}, \\ \partial_t \omega' + L'_{\bar{\omega}}[\omega'] &= -u'_i \partial^i \omega' + \partial^i \overline{u'_i \omega'} + \eta'.\end{aligned}$$

Adiabatic reduction at lowest order¹⁶:

$$\begin{aligned}\partial_t \bar{\omega} + L_{\bar{\omega}}[\bar{\omega}] &= -\partial^i \mathbb{E}_{\bar{\omega}}[\overline{u'_i \omega'}] + \bar{\eta}, \\ \partial_t \omega' + L'_{\bar{\omega}}[\omega'] &= \eta'.\end{aligned}$$

- ▶ **No UV divergences**
- ▶ **Eddy-eddy interactions** do not contribute at leading order.

The fluctuating vorticity field is an Ornstein-Uhlenbeck process characterized by the two-point correlation function $g(\mathbf{r}_1, \mathbf{r}_2, t) = \mathbb{E}_{\bar{\omega}}[\omega'(\mathbf{r}_1, t)\omega'(\mathbf{r}_2, t)]$, which satisfies the Lyapunov equation:

$$\partial_t g + L'^{(1)}_{\bar{\omega}} g + L'^{(2)}_{\bar{\omega}} g = C',$$

with $C'(\mathbf{r}_1, \mathbf{r}_2, t) = \mathbb{E}[\eta'(\mathbf{r}_1, t)\eta'(\mathbf{r}_2, t)]$ the correlation matrix of the Gaussian white noise η' .

¹⁶F. Bouchet et al. (2013).. *J. Stat. Phys.*

Adiabatic elimination of fast variables: zonal jets

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- ▶ No UV divergences
- ▶ Eddy-eddy interactions do not contribute at leading order.

Numerical simulations in the quasi-linear framework:

- ▶ Stochastic Structural Stability Theory¹⁷
- ▶ Cumulant Expansion “CE2”¹⁸

¹⁶F. Bouchet et al. (2013).. *J. Stat. Phys.*

¹⁷B. F. Farrell and P. J. Ioannou (2003).. *J. Atmos. Sci.*

¹⁸S. M. Tobias and J. B. Marston (2013).. *Phys. Rev. Lett.* J. B. Marston et al. (2016).. *Phys. Rev. Lett.*

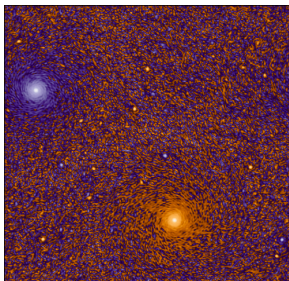
Explicit computations in the vortex condensate¹⁹

Let us go back to the periodic square box with small-scale random forcing:

$$\mathbf{v} = \mathbf{U} + \mathbf{u}, \text{ with } \mathbf{U} = U\mathbf{e}_\theta, \mathbf{u} = u\mathbf{e}_\theta + v\mathbf{e}_r \text{ and } \langle \mathbf{u} \rangle = 0,$$

$$\omega = \Omega + \omega', \text{ with } \langle \omega' \rangle = 0.$$

$$\partial_t \Omega + \mathbf{U} \cdot \nabla \Omega = -\alpha \Omega - \nabla \cdot \langle \mathbf{u} \omega' \rangle.$$



DNS: 1024^2 , $k_f = 100$, hyperviscosity, $\alpha = 1.1 \times 10^{-4}$.

¹⁹C. Herbert, A. Frishman and G. Falkovich, to appear

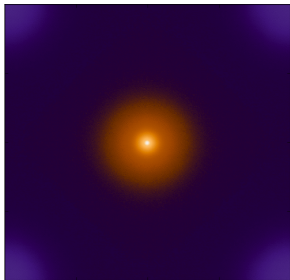
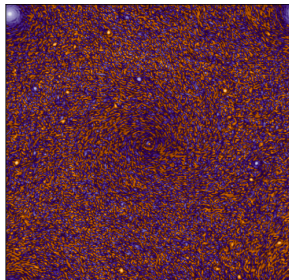
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 Ω  ω'

DNS: 1024^2 , $k_f = 100$, hyperviscosity, $\alpha = 1.1 \times 10^{-4}$.

¹⁹C. Herbert, A. Frishman and G. Falkovich, to appear

Explicit computations in the vortex condensate²⁰

Let us go back to the periodic square box with small-scale random forcing:

$$\mathbf{v} = \mathbf{U} + \mathbf{u}, \text{ with } \mathbf{U} = U\mathbf{e}_\theta, \mathbf{u} = u\mathbf{e}_\theta + v\mathbf{e}_r \text{ and } \langle \mathbf{u} \rangle = 0,$$

$$\omega = \Omega + \omega', \text{ with } \langle \omega' \rangle = 0.$$

$$\partial_t \Omega + \mathbf{U} \cdot \nabla \Omega = -\alpha \Omega - \nabla \cdot \langle \mathbf{u} \omega' \rangle.$$

Timescale separation

Perturbative expansion of the equations of motion in $\delta = \alpha L^{2/3} / \varepsilon^{1/3} \ll 1$ leads at first order to (Momentum and energy balance)¹⁹:

$$r^{-1} \partial_r (r^2 \langle uv \rangle) = -\alpha r U,$$

$$r^{-1} \partial_r (r U \langle uv \rangle) + \alpha U^2 = \varepsilon.$$

Solution:

$$U = \sqrt{3\varepsilon/\alpha}, \quad \langle uv \rangle = -r \sqrt{\alpha\varepsilon/3}.$$

Therefore $\Omega(r) = \sqrt{3\varepsilon/\alpha} r^{-1}$.

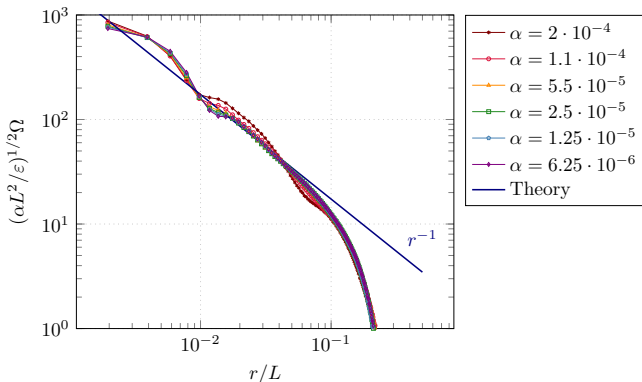
Global energy balance neglecting small-scale dissipation yields $U_{\text{rms}} = \sqrt{\varepsilon/\alpha}$.

¹⁹J. Laurie et al. (2014).. *Phys. Rev. Lett.*

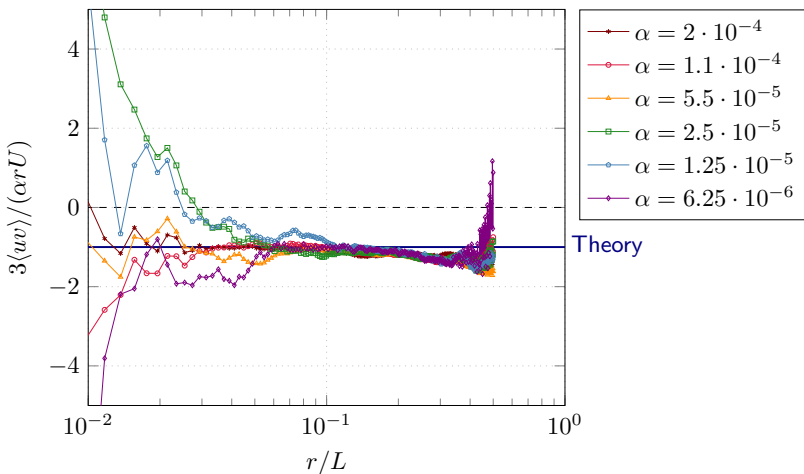
²⁰C. Herbert, A. Frishman and G. Falkovich, to appear

Explicit computation for the mean vorticity profile (DNS)²¹

Theoretical prediction: $\Omega(r) = \sqrt{3\varepsilon/\alpha}r^{-1}$



Our DNS (512² and 1024²) support the α -scaling on a wide range of α , and seem compatible with the r -scaling.

Explicit computation for the Reynolds tensor (DNS)²² $\langle uv \rangle / U^2 = O(\delta^{3/2})$ and not sign definite.DNS: 512^2 , $k_F = 100$, hyperviscosity, ~ 300000 turnover times.²²C. Herbert, A. Frishman and G. Falkovich, to appear

Summary & Prospects

Due to the existence of a small parameter, we can close asymptotically the hierarchy of moments for the 2D Navier-Stokes equations, and compute the statistics of the mean-flow (e.g. vortex condensate, jets) and fluctuations.

Salient features of the theory

- ▶ Theoretical and Numerical arguments support the timescale separation hypothesis.
- ▶ Explicit formula for the mean-flow in the vortex condensate
- ▶ Explicit computation of the average Reynolds stress tensor agrees with long time DNS.
- ▶ Dominant interactions are non-local between mean-flow and fluctuations.

Prospects

- ▶ Slow dynamics of large-scale flow (e.g. zonal jets): attractors, fluctuations, . . .
- ▶ Large deviations of the Reynolds tensor

What do we learn about mean-flow-turbulence interactions in general flows?

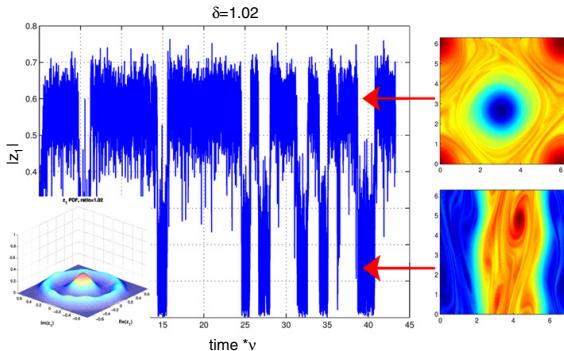
Outline

- 1 Introduction
- 2 Equilibrium Theory
- 3 Perturbative approach
- 4 Large deviations and transitions**
- 5 Conclusion

Transitions in the stochastic 2D Navier-Stokes equations

Stochastic 2D Navier-Stokes equations on a double periodic domain with aspect ratio close to one²³.

$$z_1 = \int dx dy e^{iy} \omega(x, y)$$



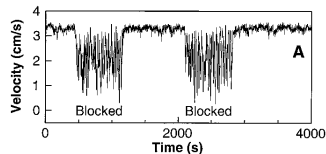
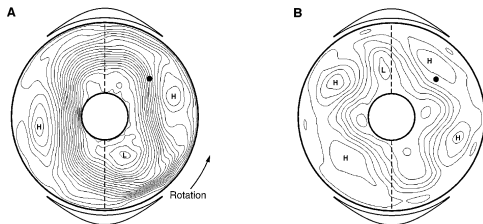
- ▶ Unidirectional flows: $|z_1| \approx 0$.
- ▶ Dipoles: $|z_1| > 0$.

Both states are close to stationary states of the Euler equations.

²³F. Bouchet and E. Simonnet (2009).. *Phys. Rev. Lett.*

Zonal-Blocked transitions

Transitions between zonal and blocked states in rotating tank experiments²⁴:



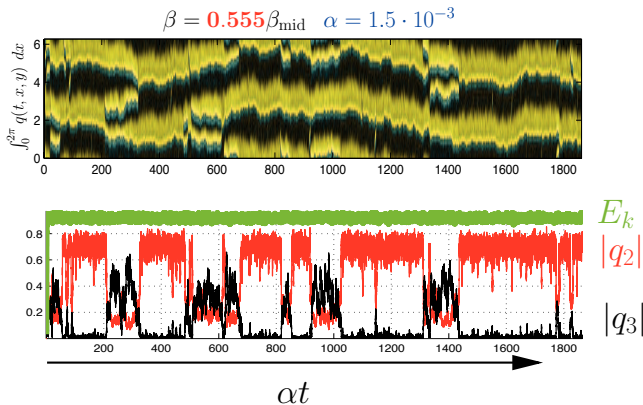
Connecting blocking and bistability is an old idea²⁵.

²⁴E. R. Weeks et al. (1997).. *Science*; Y. D. Tian et al. (2001).. *J. Fluid Mech.*

²⁵J. Charney and J DeVore (1979).. *J. Atmos. Sci.*

Rare transitions in jet dynamics

Zonal jets in the stochastic barotropic vorticity equation:

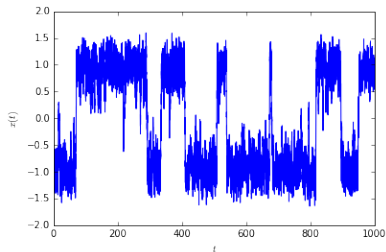
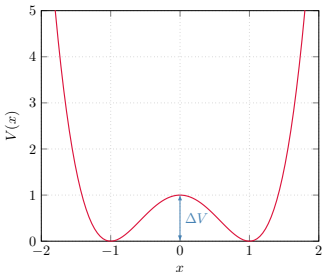


Simulations by Eric Simonnet (INLN).

Theoretical framework for noise induced transitions: the Kramers problem²⁶

Overdamped Langevin dynamics:

$$\dot{x} = -V'(x) + \sqrt{2\epsilon}\eta, \quad V(x) = (x^2 - 1)^2, \quad \mathbb{E}[\eta(t)\eta(t')] = \delta(t - t').$$

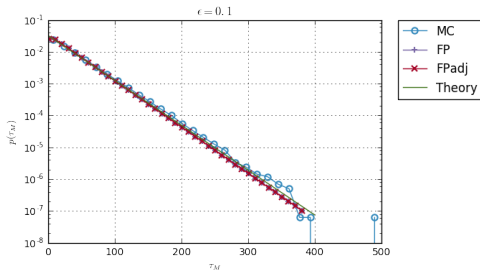
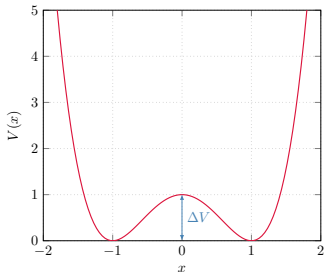


²⁶H. A. Kramers (1940).. *Physica*.

Theoretical framework for noise induced transitions: the Kramers problem²⁶

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$$\dot{x} = -V'(x) + \sqrt{2\epsilon}\eta, \quad V(x) = (x^2 - 1)^2, \quad \mathbb{E}[\eta(t)\eta(t')] = \delta(t - t').$$



Transition probability

In the weak noise limit, transition times form a Poisson point process with transition rate λ , given by

$$\lambda = \tau^{-1} e^{-\Delta V/\epsilon}$$

This is a large deviation result.

²⁶H. A. Kramers (1940).. *Physica*.

Theoretical framework for noise induced transitions: the Kramers problem²⁶

Overdamped Langevin dynamics:

$$\dot{x} = -V'(x) + \sqrt{2\epsilon}\eta, \quad V(x) = (x^2 - 1)^2, \quad \mathbb{E}[\eta(t)\eta(t')] = \delta(t - t').$$

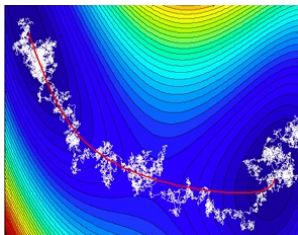
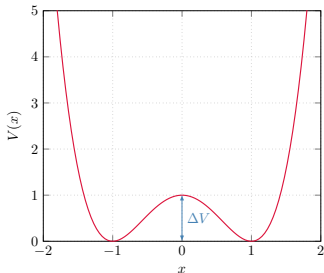


Fig. E. Vanden-Eijnden (Courant)

Instantons

Path integral formalism

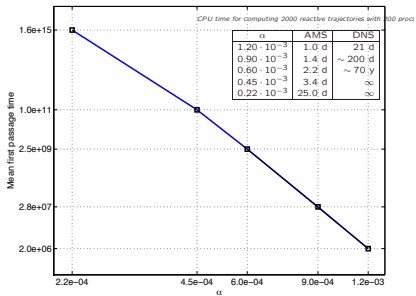
$$\mathbb{E}[\mathcal{O}] = \int \mathcal{D}[x] \mathcal{O}[x] \exp(-\mathcal{A}[x]/\epsilon), \quad \text{Action: } \mathcal{A}[x] = \frac{1}{4} \int dt (\dot{x} + V'(x))^2.$$

Instanton: most probable path: $\min_x \{\mathcal{A}[x] | x(-T) = -1, x(T) = 1\}$.

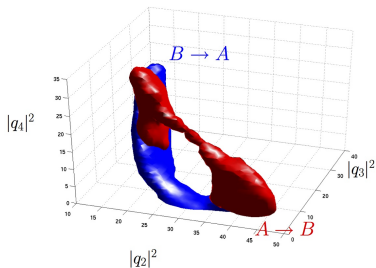
²⁶H. A. Kramers (1940).. *Physica*.

Arrhenius law and Instantons in jet transitions

Numerical algorithms to compute large deviations: dynamics biased in a controlled way²⁷.



Arrhenius law



“Instantons”

Jet transition simulations with rare event algorithm (AMS) by Eric Simonnet (INLN).

²⁷ C Gardina et al. (2011).. *J. Stat. Phys.* F. Cérou and A. Guyader (2007).. *Stoch. Anal. Appl.*

Summary and Prospects

Theoretical and numerical tools have recently been developed to study abrupt transitions in a statistical manner.

Exemples of quantities we can compute

- ▶ Probability of transition between attractors
- ▶ Most probable path (instanton theory)
- ▶ Large deviations of any observable

Recent developments

- ▶ More complex dynamics: bifurcations²⁸, non-gradient dynamics²⁹
- ▶ Large deviations and return time for time-averaged observables³⁰: applications for heat waves, cold spells, etc
- ▶ ...

²⁸C. Herbert and F. Bouchet, to appear.

²⁹F. Bouchet and J. Reygner (2016).. *Ann. Henri Poincaré*.

³⁰T. Lestang, F. Ragone, C. Herbert and F. Bouchet, to appear

Summary

Developing statistical theory for 2D and geophysical turbulence

- ▶ The mean-field theory allows one to compute statistical equilibrium states, which correspond to observed large-scale structures.
- ▶ Time scale separation allows for perturbative closure of hierarchy of moments. Explicit computation for fundamental quantities in turbulence: mean flow and Reynolds tensor.
- ▶ Abrupt transitions in turbulent flows can be studied with large deviations theory and rare event algorithms.

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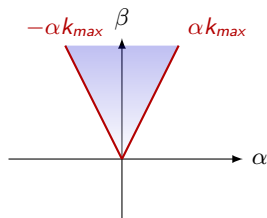
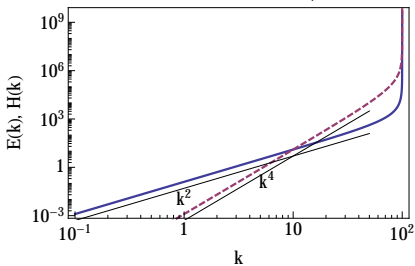
Canonical distribution for Galerkin-truncated 3D flows

Invariants: energy $E = 1/2 \int \mathbf{u}^2$ and helicity $H = \int \mathbf{u} \cdot \boldsymbol{\omega}$.
The Liouville theorem holds³¹.

Canonical probability density:

$$\begin{aligned} \rho(\{u_+(\mathbf{k}), u_-(\mathbf{k})\}) &= \frac{1}{\mathcal{Z}} e^{-\beta E - \alpha H}, \\ &= \frac{1}{\mathcal{Z}} e^{-\sum_{\mathbf{k}} [(\beta + \alpha k)|u_+(\mathbf{k})|^2 + (\beta - \alpha k)|u_-(\mathbf{k})|^2]}. \end{aligned}$$

Partition Function: $\mathcal{Z} = \prod_{\mathbf{k}} \frac{2\pi}{\sqrt{\beta^2 - \alpha^2 k^2}}$. $\beta > |\alpha| k_{\max} > 0$.



$$\begin{aligned} \langle E \rangle &= -\frac{\partial \ln \mathcal{Z}}{\partial \beta}, \\ &= \sum_{\mathbf{k}} \frac{\beta}{\beta^2 - \alpha^2 k^2}, \end{aligned}$$

$$\langle E(k) \rangle = \frac{4\pi\beta k^2}{\beta^2 - \alpha^2 k^2}.$$

*Ultraviolet Divergence*³²

³¹T. D. Lee (1952).. *Q. Appl. Math.*

³²R. H. Kraichnan (1973).. *J. Fluid Mech.*

Canonical distribution for Galerkin-truncated 2D flows

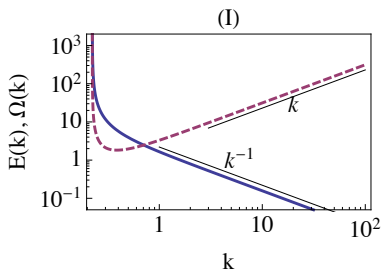
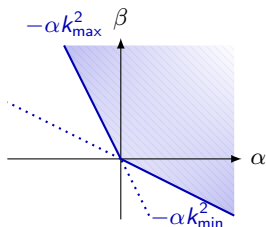
Invariants: Energy and Enstrophy:

$$\mathcal{E}[\omega] = \frac{1}{2} \int_{\mathcal{D}} \omega \psi = \frac{1}{2} \sum_{\mathbf{k} \in \mathcal{B}} \frac{|\omega_{\mathbf{k}}|^2}{k^2},$$

$$\mathcal{G}_2[\omega] = \frac{1}{2} \int_{\mathcal{D}} \omega^2 = \frac{1}{2} \sum_{\mathbf{k} \in \mathcal{B}} |\omega_{\mathbf{k}}|^2.$$

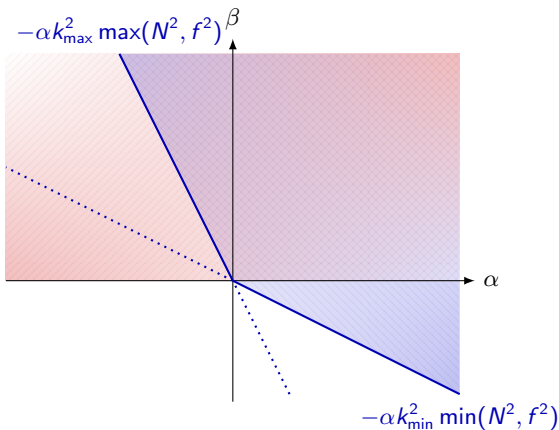
Canonical probability density³³:

$$\begin{aligned} \rho(\{\omega_{\mathbf{k}}\}_{\mathbf{k} \in \mathcal{B}}) &= \frac{1}{\mathcal{Z}} e^{-\beta \mathcal{E}[\omega] - \alpha \mathcal{G}_2[\omega]}, \\ &= \frac{1}{\mathcal{Z}} e^{-\sum_{\mathbf{k} \in \mathcal{B}} (\beta + \alpha k^2) \frac{|\omega_{\mathbf{k}}|^2}{2k^2}}, \\ \langle E(k) \rangle &= \frac{\pi k}{\beta + \alpha k^2}. \end{aligned}$$



Infrared divergence in the $\beta < 0$ regime. Inverse cascade for 2D Turbulence.

³³R. H. Kraichnan (1967).. *Phys. Fluids*; R. H. Kraichnan (1975).. *J. Fluid Mech.*



Accessible thermodynamic space for rotating-stratified flows, waves (red) and slow manifold (blue).

The helical decomposition for the 3D Euler equation

Euler equations for 3D homogeneous isotropic turbulence:

$$\begin{aligned}\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} &= -\nabla P, \\ \nabla \cdot \mathbf{u} &= 0.\end{aligned}$$

Helical decomposition in Fourier space³⁴: $\nabla \times \mathbf{h}_{\pm}(\mathbf{k}) = \pm k \mathbf{h}_{\pm}(\mathbf{k})$,

$$\begin{aligned}\mathbf{u}(\mathbf{x}) &= \sum_{\mathbf{k}} [u_+(\mathbf{k}) \mathbf{h}_+(\mathbf{k}) + u_-(\mathbf{k}) \mathbf{h}_-(\mathbf{k})] e^{i\mathbf{k} \cdot \mathbf{x}}, \\ \boldsymbol{\omega}(\mathbf{x}) = \nabla \times \mathbf{u} &= \sum_{\mathbf{k}} k [u_+(\mathbf{k}) \mathbf{h}_+(\mathbf{k}) - u_-(\mathbf{k}) \mathbf{h}_-(\mathbf{k})] e^{i\mathbf{k} \cdot \mathbf{x}}\end{aligned}$$

Automatically enforces incompressibility: $\mathbf{k} \cdot \mathbf{h}_{\pm}(\mathbf{k}) = 0$.

Energy and Helicity:

$$\begin{aligned}E &= \frac{1}{2} \int \mathbf{u}(\mathbf{x})^2 d\mathbf{x} = \frac{1}{2} \sum_{\mathbf{k}} [|u_+(\mathbf{k})|^2 + |u_-(\mathbf{k})|^2], \\ H &= \frac{1}{2} \int \mathbf{u}(\mathbf{x}) \cdot \boldsymbol{\omega}(\mathbf{x}) d\mathbf{x} = \frac{1}{2} \sum_{\mathbf{k}} k [|u_+(\mathbf{k})|^2 - |u_-(\mathbf{k})|^2].\end{aligned}$$

³⁴A Craya (1958).. *Publ. Sci. Tech. Ministère de l'Air*; J. R. Herring (1974).. *Phys. Fluids*; F. Waleffe (1992).. *Phys. Fluids A*

Macrostates and global constraints³⁵

Coarse-grained vorticity field:

$$\bar{\omega}_i = \frac{1}{n} \sum_{\alpha=1}^n \omega_{i\alpha} = \sum_{k=1}^K \sigma_k p_{ik}.$$

- ▶ The energy does not depend on the microstate but only on the macrostate

$$\begin{aligned} \mathcal{E}[\hat{\omega}] &= \frac{1}{2N^2} \sum_{(i,\alpha) \neq (j,\beta)} G_{i\alpha,j\beta} \omega_{i\alpha} \omega_{j\beta}, \\ &= \frac{1}{2M^2} \sum_{i \neq j} G_{ij} \bar{\omega}_i \bar{\omega}_j + o\left(\frac{1}{n}\right). \end{aligned}$$

- ▶ For $\hat{\omega} \in \mathfrak{M}(P)$,

$$\nu_k^T[\hat{\omega}] = \sum_{i=1}^N \nu_{ik}[\hat{\omega}] = n \sum_{i=1}^N p_{ik},$$

Global vorticity distribution constraints:

$$\frac{\nu_k^T[P]}{N} = \gamma_k.$$

³⁵C. Herbert (2015).. In: *Stochastic Equations for Complex Systems: Theoretical and Computational Topics*. Ed. by S. Heinz and H. Bessaih. Springer

The mean-field approach: thermodynamic limit

Microstates

$$\hat{\omega} = (\omega_{i\alpha})_{\substack{1 \leq i \leq M \\ 1 \leq \alpha \leq n}} \in \mathfrak{G}^N \xrightarrow{n, M, K \rightarrow +\infty} \omega(\mathbf{r}) \in L^2(\mathcal{D})$$

Macrostates

$$P = (p_{ik})_{\substack{1 \leq i \leq M \\ 1 \leq k \leq K}} \in [0, 1]^{MK} \xrightarrow{n, M, K \rightarrow +\infty} \rho(\sigma, \mathbf{r})$$

$$\forall i \in \llbracket 1, M \rrbracket, \sum_{k=1}^K p_{ik} = 1 \xrightarrow{n, M, K \rightarrow +\infty} \forall \mathbf{r} \in \mathcal{D}, \int_{\mathbb{R}} \rho(\sigma, \mathbf{r}) d\sigma = 1$$

$$\bar{\omega}_i = \frac{1}{n} \sum_{\alpha=1}^n \omega_{i\alpha} = \sum_{k=1}^K \sigma_k p_{ik} \xrightarrow{n, M, K \rightarrow +\infty} \bar{\omega}(\mathbf{r}) = \int_{\mathbb{R}} \sigma \rho(\sigma, \mathbf{r}) d\sigma$$

$$\mathcal{S}_{M, K}[P] = -\frac{1}{M} \sum_{i=1}^M \sum_{k=1}^K p_{ik} \ln p_{ik} \xrightarrow{n, M, K \rightarrow +\infty} \mathcal{S}[\rho] \equiv -\int_{\mathcal{D}} d\mathbf{r} \int_{\mathbb{R}} d\sigma \rho(\sigma, \mathbf{r}) \ln \rho(\sigma, \mathbf{r})$$

Constraints

$$\frac{1}{2} \sum_{i, j=1}^M G_{ij} \bar{\omega}_i \bar{\omega}_j = E \xrightarrow{n, M, K \rightarrow +\infty} \mathcal{E}[\rho] \equiv \frac{1}{2} \int_{\mathcal{D}^2} d\mathbf{r} d\mathbf{r}' G(\mathbf{r}, \mathbf{r}') \bar{\omega}(\mathbf{r}) \bar{\omega}(\mathbf{r}') = E$$

$$\forall k \in \llbracket 1, K \rrbracket, \frac{1}{M} \sum_{i=1}^M p_{ik} = \gamma(\sigma_k) \xrightarrow{n, M, K \rightarrow +\infty} \forall \sigma \in \mathbb{R}, \mathcal{D}_{\sigma}[\rho] \equiv \int_{\mathcal{D}} \rho(\sigma, \mathbf{r}) d\mathbf{r} = \gamma(\sigma)$$

The mean-field equation for the coarse-grained vorticity field

Mean-field equation:

$$\bar{\omega}(\mathbf{r}) = F_{\beta,\alpha}(\bar{\psi}(\mathbf{r})), \quad \text{with } F_{\beta,\alpha}(u) = -\frac{1}{\beta} \frac{d \ln \mathcal{Z}_{\beta,\alpha}(u)}{du}.$$

In particular, the equilibrium coarse-grained vorticity field is a *stationary solution of the 2D Euler equation*. Further, it is *dynamically stable*.

In general, this equation is difficult to solve:

- ▶ Nonlinear partial differential equation.
- ▶ Analytic computation of the partition function $\mathcal{Z}_{\beta,\alpha}(u)$ is rarely possible.
- ▶ Relate *a posteriori* the Lagrange parameters $\beta, \alpha(\sigma)$ to invariants $E, \gamma(\sigma)$.

Numerical methods: relaxation equations³⁶, Turkington-Whitaker algorithm³⁷, ...

When the function $F_{\beta,\alpha}$ is linear, the mean-field equation can be solved analytically. When does this happen?

- ▶ “Strong mixing” limit³⁸: $\beta \rightarrow 0$, or “low-energy” limit: $\bar{\psi} \rightarrow 0$.
- ▶ Energy-entropy variational problem
- ▶ Subclass of the full MRS equilibrium states³⁹.

Then analytical computations are possible, by introducing the eigenmodes of the Laplacian on the domain \mathcal{D} .

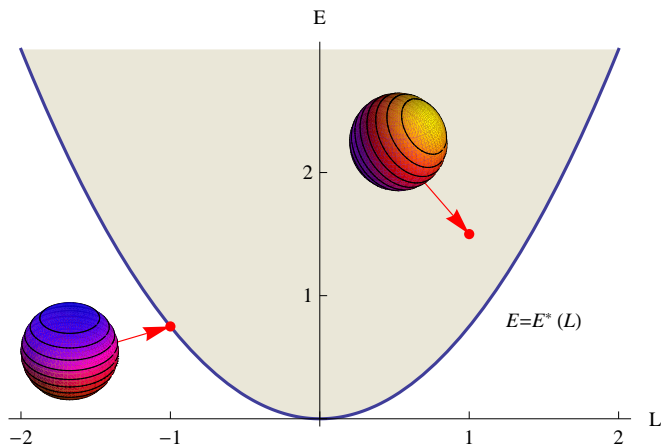
³⁶R. Robert and J. Sommeria (1992).. *Phys. Rev. Lett.* P.-H. Chavanis (2009).. *Eur. Phys. J. B*

³⁷B. Turkington and N. Whitaker (1996).. *SIAM J. Sci. Comput.*

³⁸P.-H. Chavanis and J. Sommeria (1996).. *J. Fluid Mech.*

³⁹F. Bouchet (2008).. *Physica D*

Microcanonical Phase Diagram



Second-order phase transition with spontaneous symmetry breaking.⁴⁰

⁴⁰C. Herbert et al. (2012). *J. Stat. Mech.*